

The Case for a Gaian Bottleneck: The Biology of Habitability

Aditya Chopra and Charles H. Lineweaver

Abstract

The prerequisites and ingredients for life seem to be abundantly available in the Universe. However, the Universe does not seem to be teeming with life. The most common explanation for this is a low probability for the emergence of life (an emergence bottleneck), notionally due to the intricacies of the molecular recipe. Here, we present an alternative Gaian bottleneck explanation: If life emerges on a planet, it only rarely evolves quickly enough to regulate greenhouse gases and albedo, thereby maintaining surface temperatures compatible with liquid water and habitability. Such a Gaian bottleneck suggests that (i) extinction is the cosmic default for most life that has ever emerged on the surfaces of wet rocky planets in the Universe and (ii) rocky planets need to be inhabited to remain habitable. In the Gaian bottleneck model, the maintenance of planetary habitability is a property more associated with an unusually rapid evolution of biological regulation of surface volatiles than with the luminosity and distance to the host star. Key Words: Life—Habitability—Gaia—Abiogenesis habitable zone (AHZ)—Circumstellar habitable zone (CHZ). *Astrobiology* 16, 7–22.

1. Fermi's Paradox, Hanson's Great Filter, and Bostrom's Bottlenecks

WE SEE NO EVIDENCE that our galaxy has been colonized by an advanced technological civilization. Archaeological excavations have not unearthed alien spaceships, and the optical and radio searches for extraterrestrial intelligence have not been successful (Tarter, 2001). If one assumes that once life emerges it evolves toward intelligence and technological civilizations, we are faced with Fermi's paradox: Where is everybody? [Webb (2002), Ćirković (2009); but also see Gray (2015) and Ward and Brownlee's (2000) Rare Earth hypothesis]. To put this information in context, Hanson (1998) introduced the concept of a Great Filter, describing the possible bottlenecks in the assumed progression from molecular chemistry to life, from life to intelligence, and from intelligence to galactic colonization. If the emergence of life is a rare and difficult process, then an emergence bottleneck could resolve Fermi's paradox. However, if technological civilizations inevitably destroy themselves, this self-destruction bottleneck could also resolve Fermi's paradox.

Bostrom (2008) argued that the Great Filter is a valuable tool for assessing existential risks to humanity. If the biggest barrier in the Great Filter is the emergence bottleneck, then we will find no independently evolved life on Mars, and the apparent absence of advanced technological civilizations in

our galaxy is because life's emergence is difficult and rare. In this case, humanity has already survived the biggest threat to its continued existence; the biggest bottleneck is behind us, and we can relax. However, if we find life on Mars that has evolved independently of life on Earth, this would be strong evidence that there is no emergence bottleneck. If we find such life, Bostrom argues that the biggest bottleneck—the self-destruction bottleneck—would then be in front of us. This would be “by far the worst news ever printed on a newspaper cover.” As a plausible alternative to such catastrophic logic, we introduce the concept of a Gaian bottleneck, a bottleneck that life on Earth has already passed through. If such a bottleneck exists, the discovery of *extant* independently evolved martian life might be bad news, but the discovery of *extinct* independently evolved martian life would not be.

In the standard view, the decrease in bombardment rate from ~4.5 to ~3.8 Gya is associated with making Earth more clement and thus enabling life to emerge and persist (Maher and Stevenson, 1988). In contrast to this view, we postulate a Gaian bottleneck model in which early life (on Earth and elsewhere) is not just a passive passenger but comes under strong selection pressure to actively modify and regulate its environment. The emergence of life's abilities to modify its environment and regulate initially abiotic feedback mechanisms (what we call Gaian regulation) could be the most significant factor responsible for life's persistence on Earth.

This highlights an important difference between physics-based estimates of habitability and the more unpredictable patterns of biological evolution on the highly unstable surface environments of young terrestrial planets. For example, bombardment rates inevitably decrease in the circumstellar habitable zones (CHZs) of stars, but the timescales for the evolution of Gaian regulation are probably unpredictable and would not inevitably evolve rapidly (or at all). Thus, if there is anything special about what happened on Earth to allow life to persist here, it might have less to do with the decreasing bombardment rate in the Hadean, or special chemical ingredients, or sources of free energy, or even a rare recipe for the emergence of life. The existence of life on Earth today might have more to do with the unusually rapid biological evolution of effective niche construction and Gaian regulation in the first billion years. Habitability and habitable zones would then not only be a passive abiotic property of stellar and planetary physics and chemistry (such as stellar luminosity, initial water content, and decreasing bombardment rate) but would also be a result of early life's ability to influence initially abiotic geochemical cycles and turn them into the life-mediated biogeochemical cycles that we are familiar with on the current Earth (Lenton, 1998; Lenton *et al.*, 2004; Schneider, 2004; Falkowski *et al.*, 2008; Kump *et al.*, 2009). Without rapid evolution of Gaian regulation, early extinction would be the most common fate of planetary life. Even if the emergence of life is a common feature of wet rocky planets throughout the Universe, the Gaian bottleneck model suggests that inhabited Earth-like planets would be rare.

2. If There Is an Emergence Bottleneck, It Is Probably "Recipe-Based"

2.1. No stellar bottleneck. No wet-rocky-planet bottleneck. Sun-like stars and Earth-like planets are common.

If our Sun were the most uranium-rich star in the Galaxy, and if life on Earth were uranium-based (instead of carbon-based), then we would have good reason to believe that life (either its emergence or persistence, or both) requires a rare kind of uranium-rich host star. However, we have been unable to identify any significant differences between the Sun and other stars that could plausibly be connected with an increased probability of abiogenesis. Sun-like stars are common in the Universe (Robles *et al.*, 2008). Thus, there seems to be no stellar bottleneck responsible for reducing the probability for the emergence of life.

Over the past decade, estimates for the frequency of rocky planets in, or near, the CHZ have increased (Howard *et al.*, 2012; Lineweaver and Chopra, 2012a; Bovaird and Lineweaver, 2013; Fressin *et al.*, 2013; Petigura *et al.*, 2013; Marcy *et al.*, 2014; Bovaird *et al.*, 2015; Burke *et al.*, 2015). Rocky planets in the CHZ are likely to be a common outcome of planetary formation. This result is supported by observational, theoretical, and computational models of rocky planet formation in which gas-rich protoplanetary disks evolve into dust disks in which planetesimals form and undergo oligarchic growth into planetary embryos as they differentiate into iron-nickel-rich cores, silicate mantles, and crusts dominated by incompatible lithophiles (Elkins-Tanton, 2012; Morbidelli *et al.*, 2012; Chambers, 2014; Hardy *et al.*, 2015). Models and

observations suggest that this sequence—the formation of terrestrial planets—is not a rare occurrence that needs special initial conditions. There seems to be no rocky-planet-in-the-CHZ bottleneck responsible for reducing the probability for the emergence of life.

2.2. No elemental or molecular ingredient bottleneck.

Life on Earth is made of hydrogen, oxygen, carbon, nitrogen, sulfur, and phosphorus: "HOCNSP" (Chopra *et al.*, 2010). HOCNSP are among the most abundant atoms in the Universe (Pace, 2001; Lodders *et al.*, 2009; Lineweaver and Chopra, 2012b). Since the elemental ingredients for life are the most common elements in the Universe, it is not surprising that the molecular ingredients of life are also common.

Of the elements in the Universe that form molecules, water (H₂O) is the combination of the first and second most abundant elements. Water should be a common feature of rocky planets (Raymond *et al.*, 2004, 2007; Elkins-Tanton, 2012). Radial mixing during the epoch of oligarchic growth ensures that some water-rich materials from beyond the snowline are injected into the more water-poor material of the feeding zones of rocky planets in the CHZ. Thus, it is likely that Mars and Venus both started out, like Earth, hot from accretional energy and impacts and wet from impacts of large hydrous (5–20% water) asteroids from the outer asteroid belt (Morbidelli *et al.*, 2000, 2012) and other "wet" accretionary material (Drake and Richter, 2002; Drake, 2005). Terrestrial planets in other planetary systems are also likely to start with variable, non-negligible amounts of water.

Current terrestrial life is built of monomers such as amino acids, fatty acids, sugars, and nitrogenous bases. Amino acids link together to form proteins; fatty acids link to form lipids; sugars link to form carbohydrates; and nitrogenous bases combine with sugar and phosphate to make nucleotide monomers, which link to form RNA/DNA (Lineweaver and Chopra, 2012b). Thus, life on Earth emerged when available monomers linked together to make polymers. Amino acids and other organic monomers fall from the skies in carbonaceous chondrites. We have no reason to believe that the availability of these monomers is somehow unique to Earth or the Solar System. The flux of such organics was particularly high during the first billion years of the formation of Earth, and we have no reason to believe that this will not be the case during the formation of terrestrial planets in other planetary systems. The assortment of organic compounds found in carbonaceous chondrites and the probable universality of early heavy meteoritic bombardments indicate that new planetary systems should also be supplied with organic ingredients and be conducive to the synthesis of prebiotic molecules (Ehrenfreund and Charnley, 2000; Herbst and van Dishoeck, 2009; Tielens, 2013). We expect all the ingredients of life as we know it (HOCNSP, water, amino acids, sugars, nucleic acids, HCN, and other organics) to be present and available on wet rocky planets throughout the Universe. An ingredient bottleneck seems implausible.

2.3. No free energy bottleneck since photon- and chemical redox-based energy sources are common.

Life (here and elsewhere) needs to do something for a living (Conrad and Nealson, 2001; Nealson and Rye, 2013). This living depends on extracting free energy from an

environment out of thermodynamic equilibrium (Branscomb and Russell, 2013). This extraction is based on catalyzing redox reactions or absorbing photons (Kleidon, 2012). The interiors of rocky planets throughout the Universe are denser and hotter than their surfaces. Thus, thermal gradients, density gradients, and the fluid flows they drive, set up redox potentials that can be exploited (Nisbet and Fowler, 2014). The environmental factors that enabled abiogenesis on Earth, such as the geochemical disequilibria between rocks, minerals, aqueous species, and gases, are likely to be ubiquitous on wet rocky planets throughout the Universe. In addition, for planets in the CHZ, fusion in a nearby star shines ~ 6000 K photons onto ~ 300 K surfaces and enables metabolisms based on photon capture (Annala and Annala, 2008; Lineweaver and Egan, 2008).

2.4. Recipe bottleneck? How ubiquitous are abiogenesis habitable zones?

With the absence of bottlenecks associated with stars, planets, ingredients, and sources of free energy, the only other factor that could produce an emergence bottleneck would be “recipe.” While we do not know the specifics of the prebiotic chemistry and geochemical environments necessary for life to emerge (*e.g.*, Orgel, 1998; Stüeken *et al.*, 2013), the view that life will emerge with high probability on Earth-like planets is shared by many scientists. Darwin (1871) speculated that the environment, ingredients, and energy sources needed for the origin of life could be common, when he wrote about life emerging in “some warm little pond with all sort of ammonia and phosphoric salts, light, heat, electricity present, that a protein compound was chemically formed, ready to undergo still more complex changes.”

de Duve (1995) has been a more recent proponent of the view that the emergence of life is a cosmic imperative: “Life is either a reproducible, almost commonplace manifestation of matter, given certain conditions, or a miracle. Too many steps are involved to allow for something in between.” Discarding miracles, de Duve leaves us with life as an “almost commonplace manifestation of matter.” As far as we know, Darwin’s warm little pond or de Duve’s “certain conditions” are common on Earth-like planets and may be sufficient for the emergence of life. Biota throughout the Universe would emerge from chemistry through protobiotic molecular evolution (*e.g.*, Eigen and Winkler, 1992; Eschenmoser and Volkan Kisakürek, 1996; Orgel, 1998; Ward and Brownlee, 2000; Martin and Russell, 2007; Russell *et al.*, 2013).

There is some tension between this conclusion and the inability of synthetic biologists to produce life, despite having access to a wide variety of ingredients, environmental setups, and energy sources. Plausible explanations for this tension include (1) there is an emergence bottleneck due to a convoluted recipe whose requirements only rarely occur naturally; or (2) there is no emergence bottleneck—the recipe for life is simple—but we are not as imaginative or as resourceful as nature, so we have not replicated the recipe in the relatively short time we have been investigating the origin of life. We conclude that the idea of an emergence bottleneck is still plausible. However, the evidence for it is getting weaker as we find out more about the proto-biotic molecular evolution that led to the emergence of life on Earth (*e.g.*, Benner, 2013;

England, 2013; Nisbet and Fowler, 2014; Sousa and Martin, 2014). This weakness provides motivation for alternative explanations for the apparent paucity of life in the Universe. The Gaian bottleneck hypothesis is one such alternative.

As we learn more about the origin of life, we can begin to define an abiogenesis habitable zone (AHZ) where the requirements for life’s emergence are met. Many initially wet rocky planets in the CHZ may possess the necessary and sufficient conditions to get life started (AHZ in Fig. 1). The habitability requirements for the origin of life may be substantially different from, and more specific than, the requirements to maintain life on a planet (as in the difference between the need to have a spark plug to start an engine and a radiator to keep it from overheating). In Fig. 1, we have assumed that the AHZ is not necessarily a subset of currently inhabited planetary environments.

3. Emergence Bottlenecks versus Gaian Bottlenecks

An emergence bottleneck is illustrated in Fig. 2. The left panel shows a hypothetical planet with non-evolving planetary conditions. The right panel shows a more plausible planet that initially had some habitable regions but, through volatile evolution or other transient factors, lost its surface water and evolved away from habitable conditions (*e.g.*, a runaway greenhouse or runaway glaciated planet). Without significant abiotic negative feedback mechanisms, the surface environments of initially wet rocky planets are volatile and change rapidly without any tendency to maintain the habitability that they may have temporarily possessed as their early unstable surface temperatures transited through habitable conditions (Fig. 6C).

If there is no emergence bottleneck (Fig. 3), typical wet rocky planets have initial conditions compatible with the emergence of life (AHZ). We postulate that almost all initially wet rocky planets on which life emerges (left panel of Fig. 3) quickly evolve like the abiotic planets represented in the right panel of Fig. 2. This unregulated evolution of planetary environments away from habitable conditions constrains the duration of life’s existence on the planet. We call this early extinction of almost all life that ever emerges the *Gaian bottleneck*. In rare cases (for example on Earth), life will be able to evolve quickly enough to begin to regulate surface volatiles through the modification of abiotic feedbacks (right panel of Fig. 3). The potentially relevant feedbacks involved in such early Gaian regulation are illustrated in Fig. 5 and discussed in Table 1 and Section 5.

4. Early Extinction from Impacts and Devolatilization

4.1. Bombardment, impact frustration, and extinction

During the late phases of Earth’s accretion (~ 4.5 to ~ 4.0 Gya), episodes of cold “Norse ice-hell” were punctuated by brief periods of hot inferno with magma oceans (Nisbet, 2002; Elkins-Tanton, 2008). Frequent, large, random impacts produced wide-ranging, unstable temperatures. The largest impacts were so large that any life that did emerge during the transitory habitable periods was probably bombarded into extinction. These early bombardment-induced extinctions have been called the impact frustration of life (Maher and Stevenson, 1988; Sleep *et al.*, 1989, 2014; Davies and Lineweaver, 2005; Marchi *et al.*, 2014).

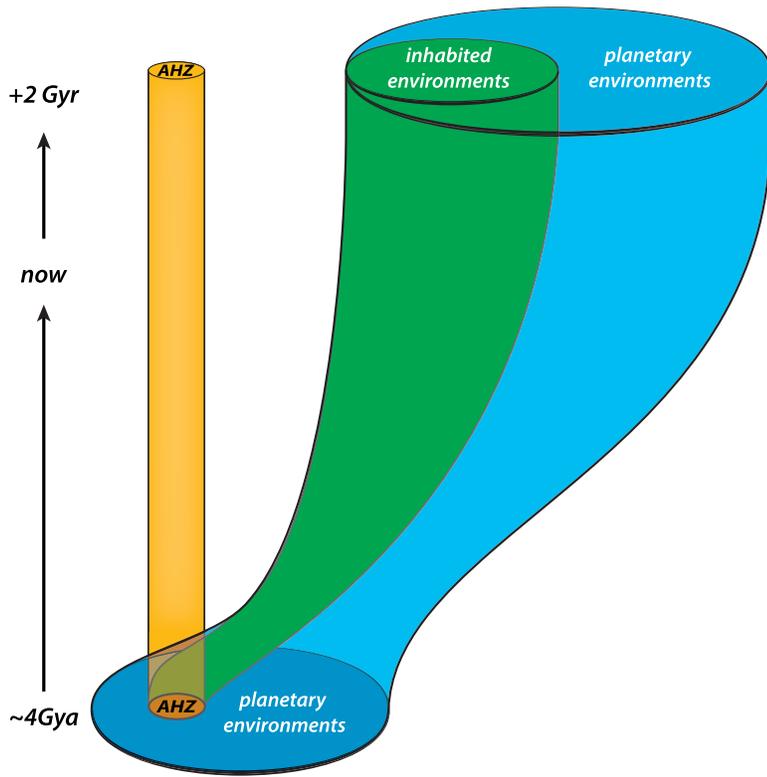


FIG. 1. The conditions needed for the origin of life are in the abiogenesis habitable zone, or AHZ. Inhabited environments (green) are a subset of planetary environments (blue). Both of these can change with time. The AHZ conditions are probably narrower than the broader conditions to which life can now adapt (“inhabited environments”). Through its management of the greenhouse and its partitioning of reductants and oxidants, the activity of life increases the range of inhabited environments (Nisbet *et al.*, 2007). Hence, the green cylinder emerges out of the AHZ and gets broader with time. More specific reasons for this broadening include (1) the evolution of increasingly efficient catalytic enzymes offering tighter control over reaction rates; (2) the ability of new enzymes to access the energy from different redox pairs providing larger $|\Delta G|$ values (Nealson and Conrad, 1999); and (3) the evolution of ecosystems (Smith and Morowitz, 2016), global-level niche construction, and global biogeochemical feedback cycles (see Section 5), which we refer to as the “evolution of Gaian regulation of the biosphere.” (Color graphics available at www.liebertonline.com/ast)

Emergence Bottleneck

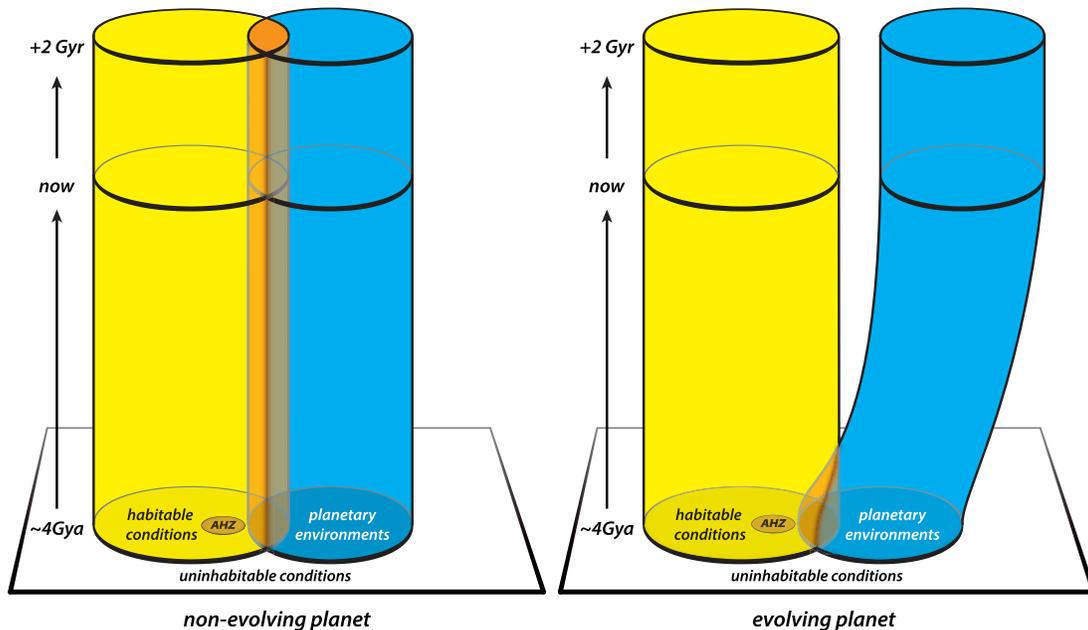


FIG. 2. Emergence Bottleneck: planets on which life does not get started. Here, we show habitable conditions (yellow) and planetary environments (blue), from the time of planetary formation at the bottom to ~ 6 billion years later at the top. Life will not emerge on either of these planets since their initial planetary environments do not overlap with the AHZ where life could get started. Both of these planets are initially habitable since their environments overlap with habitable conditions. *Left panel:* In this unrealistic, non-evolving model, planetary environments do not change with time. Habitable regions remain uninhabited because life does not get started. Such planets are uninhabited but remain habitable. *Right panel:* Parts of this evolving planet were habitable, but life does not emerge. The planet undergoes abiotic evolution and moves quickly away from habitability. We argue (Section 4.2) that evolution away from habitability is probably the default for initially wet rocky planets. We would find such planets to be uninhabited (and if older than ~ 1 billion years, uninhabitable)—consistent with the idea that a planet has to be inhabited to remain habitable. (Color graphics available at www.liebertonline.com/ast)

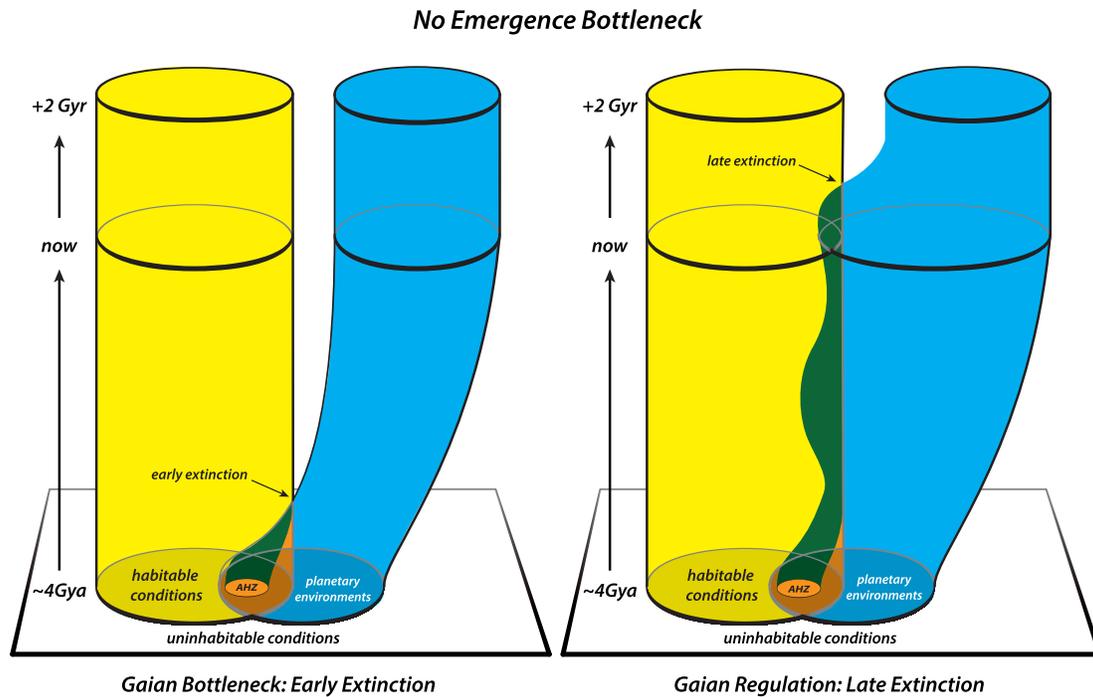


FIG. 3. The AHZ may be a very common subset of the environments of rocky planets that are wet during their first billion years. As in the left panel of Fig. 2, we assume that these planets are unregulated by any abiotic negative feedbacks and have no tendency to maintain habitability. We assume that life gets started on both of these planets, so there is no “emergence bottleneck” (Section 2) as there was in Fig. 2. *Left panel:* Life is unable to evolve rapidly enough to control runaway positive feedbacks. Gaian regulation does not emerge fast enough to maintain habitability. Thus, we have a “Gaian bottleneck.” We propose that most wet rocky planets are of this kind. *Right panel:* In rare cases (as on Earth), Gaian regulation evolves fast enough to make it through the Gaian bottleneck and keep at least part of the planet habitable for billions of years. Biospheric regulation maintains the habitability of the planet until ~5 billion years after formation, at which time, increasing stellar luminosity and loss of water cause life to go extinct. Extinction happens in both panels but much later in the right panel. This figure illustrates our hypothesis that the emergence and rapid extinction of life may be quite common (left) but that the emergence of life followed by the evolution of Gaian regulation and the long-term persistence of life could be quite rare (right). (Color graphics available at www.liebertonline.com/ast)

The early heavy bombardment of Earth decreased by more than ~13 orders of magnitude during this period (Bland, 2005; Köberl, 2006; Fig. 6A). As the rate of planetary accretion slowed, the impact rate and the size of the largest impactors decreased.

Habitable conditions became, at least fleetingly, more available for life (Abramov and Mojzsis, 2009; Abramov *et al.*, 2013). Impact heating may have provided strong selection pressure on early life to evolve into deep environments to survive thermal perturbations (Sleep *et al.*, 1989; Nisbet and Sleep, 2001; Mat *et al.*, 2008).

We have no reason to believe that these processes are specific to Earth or to our solar system. The surfaces of Earth-sized rocky planets will undergo an early heavy bombardment that produces severe intermittent temperature pulses. A decreasing bombardment rate is likely to be a generic process that controls the emergence and early persistence of life on wet rocky planets near the CHZs of host stars throughout the Universe. As the bombardment rate decreases, these rocky planets cool down and can harbor, at least temporarily, habitable environments where life can emerge from primordial soups, hydrothermal vents, or any other AHZ candidate. Whether life usually persists after this emergence is what we are calling into question.

We postulate that the combination of heavy bombardment, volatile evolution, and thermal instability almost always conspires to eliminate incipient life before it has a chance to evolve sufficiently to regulate initially abiotic global cycles. We also postulate that, exceptionally, life on Earth was able to counter the effects of volatile evolution, thermal instability, and the general abiotic tendency to drift away from habitability. In other words, we argue that the same catastrophic events that life on Earth seems to have overcome usually lead to extinction. In the first billion years on a wet rocky planet in the Universe, impacts and an inability to control surface environments are usually not just frustrating, but fatal.

4.2. Water loss and tendency to evolve away from habitability

Liquid water is not easy to maintain on a planetary surface. The initial inventory and the timescale with which water is lost to space due to a runaway greenhouse, or frozen due to ice-albedo positive feedback, are poorly quantified, but plausible estimates of future trajectories have been made. On Earth, dissociation of water vapor by UV radiation in the upper atmosphere is ongoing and will eventually (~1–2 billion years from now) lead to the loss of water from

the bioshell and the subsequent extinction of life on Earth (Caldeira and Kasting, 1992; Franck, 2000; Lenton and von Bloh, 2001; Franck *et al.*, 2002; von Bloh *et al.*, 2005).

In our solar system, Venus, Earth, and Mars are usually assumed to have started out in similar conditions: hot from accretion and wet from the impacts of aqueous bodies from beyond the snowline. However, atmospheric evolution of these planets diverged significantly (Kasting, 1988; Kulikov *et al.*, 2007; Driscoll and Bercovici, 2013). The simplest interpretation of the D/H ratios of Venus:Mars:Earth:Sun (2000:70:10:1) is that (i) Venus has lost the vast majority of its H₂O; (ii) Mars lost about 85% of its initial water content, and the rest froze into the polar ice-caps and subsurface permafrost (Kurokawa *et al.*, 2014; Villanueva *et al.*, 2015); and (iii) Earth was able to keep a larger fraction of its H₂O (Pope *et al.*, 2012; Fig. 4).

However, the answer to the question “Why didn’t Earth undergo a runaway greenhouse like Venus or a runaway glaciation like Mars?” may have as much, or more, to do with life on Earth than with Earth’s distance from the Sun. The biotic mechanisms of how this preservation has been achieved have been discussed in the context of the Gaia hypothesis by Harding and Margulis (2010). The early devolatilization of Earth-like planets around M stars due to an extended pre-main sequence period of high extreme UV flux (above the dissociation energy of water, ~5 eV) could apply to some extent to Earth-like planets around more massive stars (Luger and Barnes, 2015; Tian, 2015).

The amount of water (and volatiles in general) deposited or devolatilized during the late accretion phase of rocky planet formation in the Universe is highly variable (Raymond *et al.*, 2004, 2009) and can produce desert worlds (Abe *et al.*, 2011), ocean worlds (Léger *et al.*, 2003), and probably everything in between. Abiotic volatile evolution will be rapid, stochastic, and hostage to the timing, mass, volatile content, and impact parameters of the largest impactors and the runaway feedbacks they could induce.

We argue that abiotic habitable zones are available initially and fleetingly to wet planets within a wide range of orbital radii (~0.5 to ~2 AU) because of the thermal instability of their surfaces. Wide-ranging unstable temperatures could provide transitory abiotic habitable zones during

the first half billion years after formation (Nisbet, 2002; Fig. 6B, 6C). There are two ways to influence the surface temperature of a planet (Fig. 5): change the albedo (gray loops) or change the greenhouse gas content of the atmosphere (blue loops) (Kasting, 2012). The amount and the phases of the volatiles (H₂O, CO₂, CH₄) of rocky planet atmospheres control both the albedo and greenhouse warming. Albedo and greenhouse warming, in turn, control the amount and phases of the volatiles. Strong positive feedback cycles (left side of Fig. 5) may lead to both (i) runaway greenhouse (temperatures too hot for life) with runaway loss of atmosphere (hydrogen loss and thus water loss) or (ii) runaway glaciation (lowering the temperature and/or water activity to levels not conducive to life).

4.3. The implausibility of the carbonate-silicate cycle (and other abiotic negative feedback) in the first billion years

In his original estimate of the continuous habitable zones, Hart (1979) considered runaway greenhouse and runaway glaciation feedback but did not account for the negative feedback of silicate weathering on his models. The resulting continuous habitable zone of 0.95–1.01 AU had such a narrow width that he wrote: “It appears, therefore, that there are probably fewer planets in our galaxy suitable for the evolution of advanced civilizations than has previously been thought.”

Such a narrow CHZ could help solve Fermi’s paradox (*e.g.*, Webb, 2002, Solution 36, p. 158). More recent work has taken into account a stabilizing negative feedback loop associated with the recycling of CO₂ by plate tectonics. Walker *et al.* (1981) proposed a greenhouse gas–based negative feedback process employing the carbonate-silicate cycle through the mechanism of silicate weathering. Increasing surface temperatures increases the silicate weathering rate, freeing up more Ca²⁺ and other cations that combine with CO₂ (via aqueous bicarbonate HCO₃⁻) to produce insoluble carbonates. Thus, when the temperature rises, more atmospheric CO₂ is sequestered, and temperature decreases (Walker, 1985, and the blue negative feedback loop on the right side of Fig. 5). This abiotic negative feedback is largely responsible for moving the outer edge of the CHZ from 1.01 AU, as estimated by Hart (1979), to

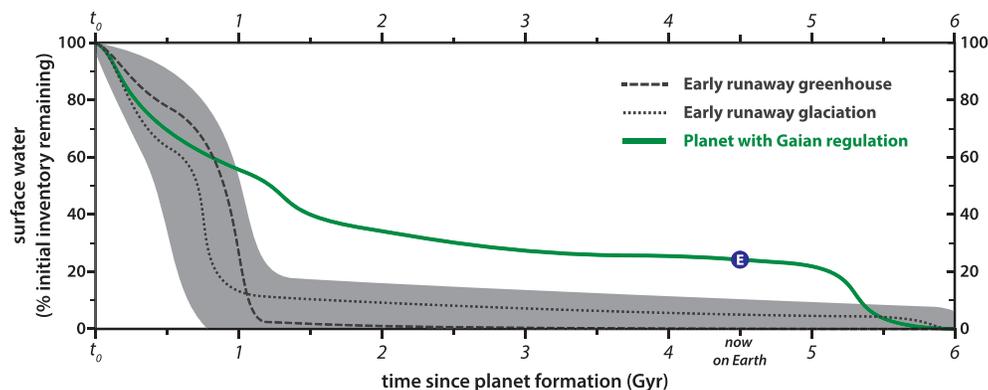


FIG. 4. Schematic illustration of the water loss caused by impacts and hydrogen escape. Hydrogen escape may entirely desiccate a rocky planet within a few billion years (Lovelock, 2005). Desiccation is inevitable as the host star luminosity increases, a cold trap is lost, and the stratosphere becomes moist (Lenton and von Bloh, 2001; O’Malley-James *et al.*, 2015). (Color graphics available at www.liebertonline.com/ast)

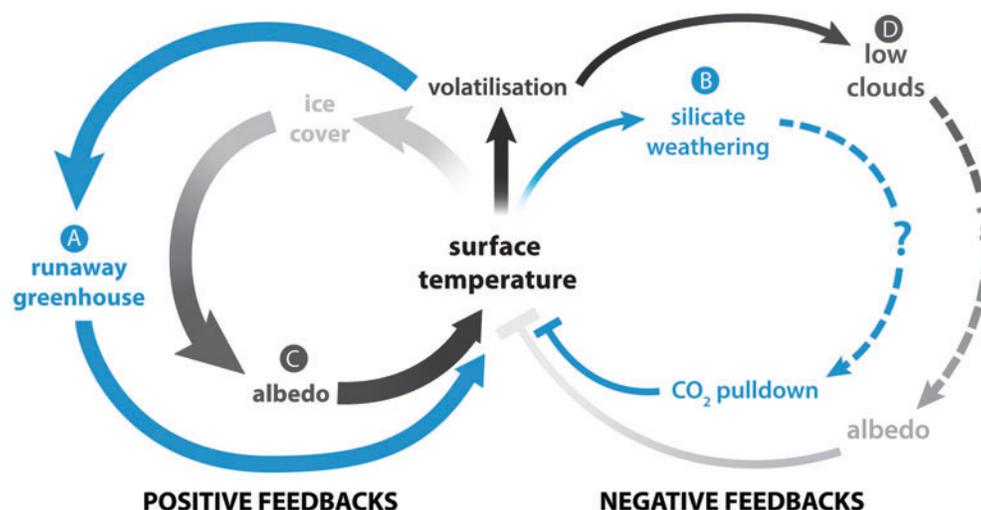


FIG. 5. Early Abiotic Feedbacks. During the first billion years after the formation of Earth (or of Earth-like planets), abiotic positive feedbacks (left) can lead to runaway surface temperatures outside the habitable range (both too hot and too cold). These positive feedbacks lead to the loss of liquid water [either from hydrogen escape to space or condensation into ice (Fig. 4)]. Abiotic negative feedbacks (right) have been invoked to stabilize surface temperatures, but they may not be significant in the first billion years, hence the dashed lines and the question marks (Section 4.3). As life evolves, it can strengthen or weaken these initially abiotic geochemical feedback loops and turn them into biogeochemical cycles and feedback loops. Evolving life can insert itself into these feedbacks at the points labeled A, B, C, and D (Table 1 and Section 5). (Color graphics available at www.liebertonline.com/ast)

TABLE 1. ABIOTIC FEEDBACK PROCESSES ACTIVE DURING THE FIRST BILLION YEARS OF A WET ROCKY PLANET’S HISTORY AND POTENTIAL BIOTIC ENHANCEMENTS OF THE FEEDBACK CYCLE

Feedback	Mechanism	Feedback type	Potential biological mediation
Greenhouse	Greenhouse gases (Ingersoll, 1969; Abe <i>et al.</i> , 2011)	Positive	A (Catling <i>et al.</i> , 2001; Kasting, 2012; Goldblatt <i>et al.</i> , 2009; Harding and Margulis, 2010)
Greenhouse	Silicate weathering (Walker <i>et al.</i> , 1981)	Negative?	B (Lovelock and Whitfield, 1982; Schwartzman and Volk, 1989; Catling <i>et al.</i> , 2001; Rosing <i>et al.</i> , 2006; Höning <i>et al.</i> , 2014)
Albedo	Ice albedo (Budyko, 1969; Hoffman, 1998; Kopp <i>et al.</i> , 2005)	Positive	C (Harding and Margulis, 2010; Watson and Lovelock, 1983)
Albedo	Low clouds (Abe <i>et al.</i> , 2011)	Negative?	D (Rosing <i>et al.</i> , 2010)

the more modern, larger values of 1.5–1.7 AU (*e.g.*, Kasting *et al.*, 1993; Kopparapu *et al.*, 2013).

Since the temperature-dependent carbonate-silicate cycle provides a negative feedback, it could have been responsible for the long-term stabilization of Earth’s surface temperature. With a sufficiently high silicate weathering rate, even a lifeless planet could remain habitable. However, since temperature-dependent silicate weathering requires sub-aerial weathering of silicate rocks (either granitic or basaltic), the magnitude of the negative feedback of silicate weathering is roughly proportional to the amount of sub-aerial continental crust. In the first billion years of Earth’s history, the fraction of the surface of Earth where sub-aerial erosion would have been possible may have been extremely small (Flament *et al.*, 2008; Abbot *et al.*, 2012; Dhuime *et al.*, 2015). Flament *et al.* (2008) modeled the sub-aerial weathering as a function of time and estimated that in the late Archean (~2.5 Gya) 2–3% of the Earth’s surface was sub-aerial continent.

The little continental crust present was largely submerged. Thus, it is likely that early in Earth’s history the negative feedback of the carbonate-silicate cycle may have been inop-

erative or at least significantly less effective than today. This undermines the main negative feedback mechanism proposed to stabilize surface temperature on wet rocky planets for the first billion years or so, when they are most likely to experience runaway greenhouse or runaway glaciation due to high inventories of primordial greenhouse gases, higher bombardment rates, and higher volcanism; hence, the “?” associated with this negative feedback loop in Fig. 5 and Table 1. Without this abiotic feedback cycle to extend the outer edge of the CHZ, the much narrower Hart-like continuously habitable zone (Fig. 6B) becomes more plausible.

The other abiotic negative feedback on surface temperature shown in Fig. 5 is associated with low clouds: higher temperatures → more volatilization → more low-altitude clouds → higher albedo → lower temperatures. Low clouds increase albedo and decrease surface temperatures more than they contribute toward raising the surface temperature because of the greenhouse effect associated with clouds (Abe *et al.*, 2011). This is problematic because increasing volatilization produces both more low-altitude clouds, which could cool Earth due to their higher albedo, and more high-altitude clouds, which could

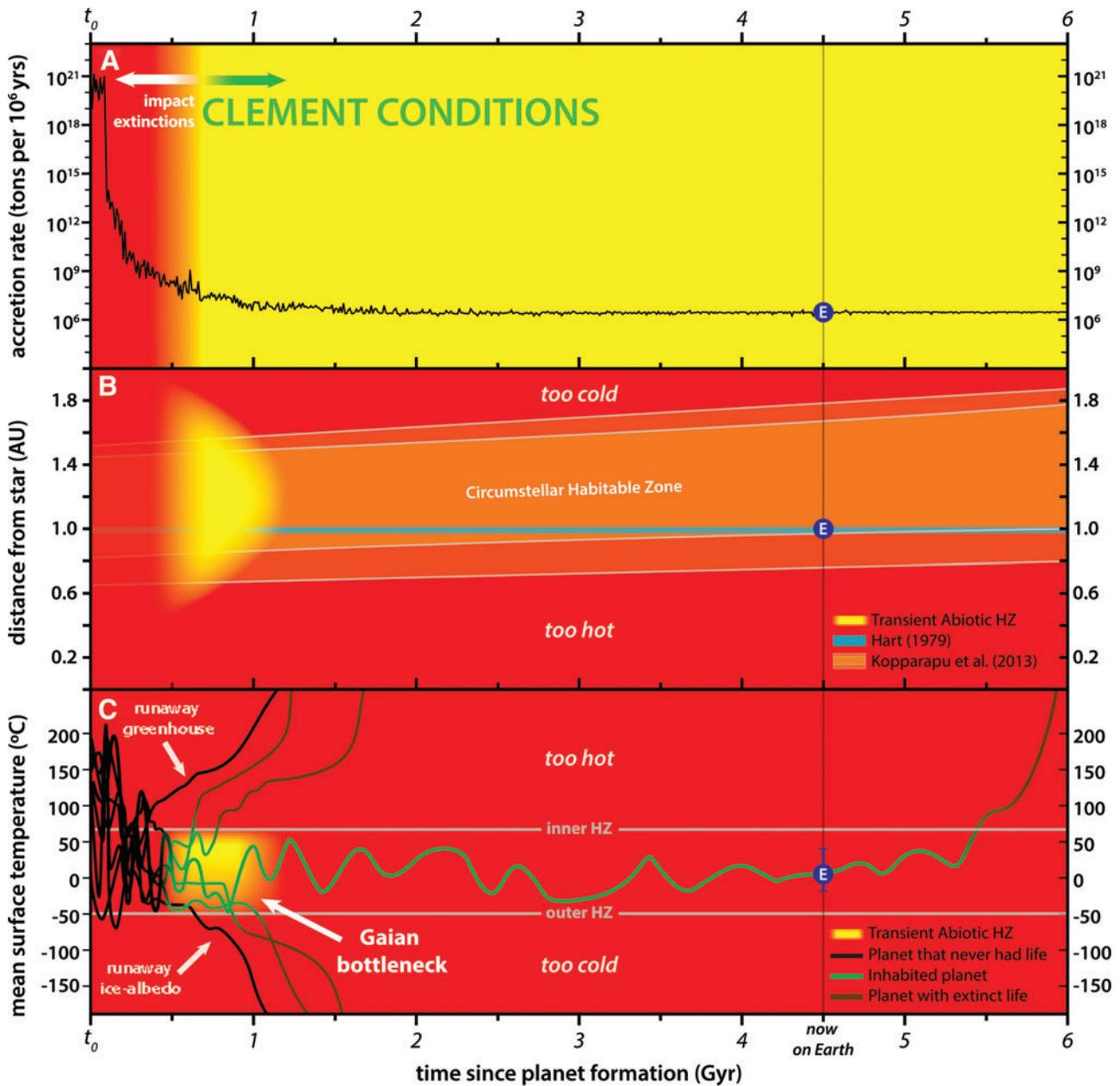


FIG. 6. Bombardment, habitable zones, and the Gaian bottleneck. (A) Early heavy bombardment precludes life for the first ~ 0.5 billion years, indicated by red in all three panels. These impacts produce heat pulses and both deliver and remove volatiles (Elkins-Tanton, 2011). Thus, the amount of H_2O at the surface is highly variable during this period. The phase of the H_2O at the surface is also highly variable during this period because of impact-induced alternation between greenhouse warming and ice-albedo runaway. (B) The width of the CHZ is usually considered to be a function of physics and chemistry but is often computed without the largely uncertain influence of clouds and placed between Venus (0.7 AU) and Mars (1.5 AU). Here the blue region is from the work of Hart (1979), and the orange regions are based on estimates by Kopparapu *et al.* (2013) for the conservative (light orange) and optimistic (dark orange) limits. Life plays no role in these computations. The $\sim 30\%$ increase in the luminosity of the Sun since its formation is responsible for the outward migration of the traditional CHZ and some of the narrowness of Hart's continuously habitable zone. The yellow zone in (B) represents our speculative version of a short-lived abiotic habitable zone at the tail end of the impact-induced thermal instabilities shown in the first billion years of panel (C). The width of the abiotic habitable zone begins with a fairly wide range of semimajor axes but lasts only from ~ 0.5 to ~ 1 Gyr and then shrinks to zero. After ~ 1 Gyr, rapid impact-induced thermal excursions diminish, and surface temperatures drift away and run away from habitability. Planets become devolatilized because of the runaway greenhouse effect (top of panel C), and because liquid water condenses out into ice due to the runaway ice-albedo effect (bottom of panel C), with no abiotically stable zone between them. The early evolution of Gaian regulation may be the main feature responsible for maintaining the surface temperature of Earth within a habitable range for the past ~ 4 billion years (Lovelock, 2000). (Color graphics available at www.liebertonline.com/ast)

have a stronger greenhouse effect than albedo effect and thus increase surface temperatures (Goldblatt and Zahnle, 2011; Leconte *et al.*, 2013). For this reason, the effects of clouds are often considered the biggest source of uncertainty in global climate models and thus the “?” associated with this cycle in Fig. 5 and Table 1.

While it may be possible to vary albedo and greenhouse gases within some plausible range and construct a wide CHZ, without negative feedback, there is no justification for tuning these abiotic variables to maintain habitability. We postulate that the abiotic stabilizing feedbacks (two cycles on the right side of Fig. 5) were probably negligible on early Earth. In their absence, it is hard to understand how habitability would have been maintained. Driverless cars do not stay on roads. Without significant abiotic stabilization, we propose that the most plausible default becomes the abiotic tendency to evolve away from habitability shown in Figs. 1 and 6C.

Just because Earth is at 1 AU and has been inhabited for ~4 billion years does not mean that there is a physics-based, biology-independent, computable continuous habitable zone. With thermal instability and increasing stellar luminosity, it is not clear that a physics-based continuously habitable zone even exists. There may be no range of orbital distances (or any region of multidimensional abiotic parameter space) for which the surface environments of initially wet rocky planets have sufficiently strong abiotic negative feedback to maintain habitability. If this is the case, purely abiotic computations of a continuously habitable zone may be misleading, and Gaian regulation becomes a plausible explanation for the continuously inhabited habitable zone in which we find ourselves.

5. The Biology of Habitability in the First Billion Years

It is usually assumed that the CHZ is determined by abiotic physical parameters: stellar mass and luminosity, planetary mass and atmospheric greenhouse gas composition, surface albedo, and sometimes clouds. More recently planetary spin, orbital eccentricity, obliquity, and initial water content have been added to the list of physical parameters (*e.g.*, Gaidos *et al.*, 2005; Gonzalez, 2005; Lammer *et al.*, 2009; Güdel *et al.*, 2014; Shields *et al.*, 2016). Here, we argue that these abiotic parameters can fleetingly enable the emergence of life but cannot maintain habitable surface conditions with liquid water. As the early heavy bombardment subsides, strong selection pressure on life begins to regulate, control, and even dominate the mechanisms that create or maintain the temperatures and pressures at the surface of a planet that allow liquid water. If so, then biology (rather than physics or chemistry) can play the most important role in maintaining habitability.

In addition to the abiotic environmental changes (due to bombardment and devolatilization), there could be a long struggle that starts early between life and an environment that does not, abiotically, stay habitable. Feedback between life and environment may play the dominant role in maintaining the habitability of the few rocky planets in which life has been able to evolve Gaian regulation quickly. If life gets started on a planet, there are many potential ways in which life can regulate the mechanisms that create or maintain the temperatures and pressures needed for liquid water (Schneider and Boston, 1991; Schneider, 2004; Harding and Margulis, 2010). Gaia researchers propose that life on Earth evolved to become integrated into previously abiotic feedback systems

that can modify or regulate surface temperature and the hydrological cycle (*e.g.*, Lenton, 1998; Nisbet *et al.*, 2012). Life can evolve to enhance and regulate the feedback loops (biological mediation processes A–D in Fig. 5 and Table 1). Biologically mediated feedback loops are stabilizing or Gaian (Ricklefs and Miller, 2000). For habitability to be maintained, life could down-regulate the positive runaway feedback loops and enhance the negative feedback loops. On Earth, life began to modulate the greenhouse gas composition of the atmosphere as soon as life became widespread (Nisbet, 2002; Nisbet *et al.*, 2012; Nisbet and Fowler, 2014; Johnson and Goldblatt, 2015).

The use of the Gaia hypothesis in ecology was reviewed by Free and Barton (2007). They argued [in agreement with Dawkins (1982)] that selection for global stability is implausible. However, they defined a Probable Homeostatic Gaia model: a planet “with appropriate starting conditions for life will probably generate a biosphere the lifespan of which will be extended, rather than reduced, by life-environment feedback.” In their Probable Homeostatic Gaia model, a “network of life-environment interactions, largely dependent on the by-product effects of evolved traits, leads to global stability.” This biology-based global stability is what we call Gaian regulation. We are invoking its rapid evolution to stabilize the early volatile and thermal instabilities of wet rocky planets.

If Gaian regulation plays the dominant role in maintaining liquid water at the surface (Harding and Margulis, 2010), then the width of the CHZ would depend more on the quirks of biological evolution than on the more deterministic physics and chemistry that can be more easily modeled. For most rocky planets in the CHZ to remain habitable, they may have to be inhabited: “habitability depends on inhabitation and the width of the habitable zone is difficult to characterize” (Goldblatt, 2015).

In Gaian literature, it is usually proposed that Earth became a Gaian planet in the Proterozoic (~2.5 Gya) and has been one ever since (Harding and Margulis, 2010, p. 45). We are proposing that the onset of Gaian regulation could have occurred more than a billion years earlier, for example, through the production, consumption, and regulation of greenhouse active gases such as H₂, CO₂, and CH₄. If there is a biotic solution for the faint early Sun paradox (Sagan and Mullen, 1972; Sagan and Chyba, 1997; Feulner, 2012) based on higher concentrations of CO₂, on biotic methanogenesis, or on biotic albedo regulation, then this solution would necessarily have evolved quickly during the transient abiotic habitable zone (yellow regions in Fig. 6B, 6C) (Walker, 1985; Pavlov *et al.*, 2003; Haqq-Misra *et al.*, 2008; Rosing *et al.*, 2010).

6. Predictions and Tests of the Gaian Bottleneck Hypothesis

After the early heavy bombardment, life emerges with some probability on initially wet rocky planets in the CHZ. However, due to large impacts and unstable abiotic volatile evolution with no tendency to maintain habitability, almost all life goes extinct early—with the rare exceptions of life that has undergone unusually rapid evolution and obtained some level of Gaian regulation. The most significant predictions of this Gaian bottleneck model can be seen in Fig. 7 by comparing panels A and B with C. In panels A and B, wherever life emerges, it persists for billions of years. Thus, it has time

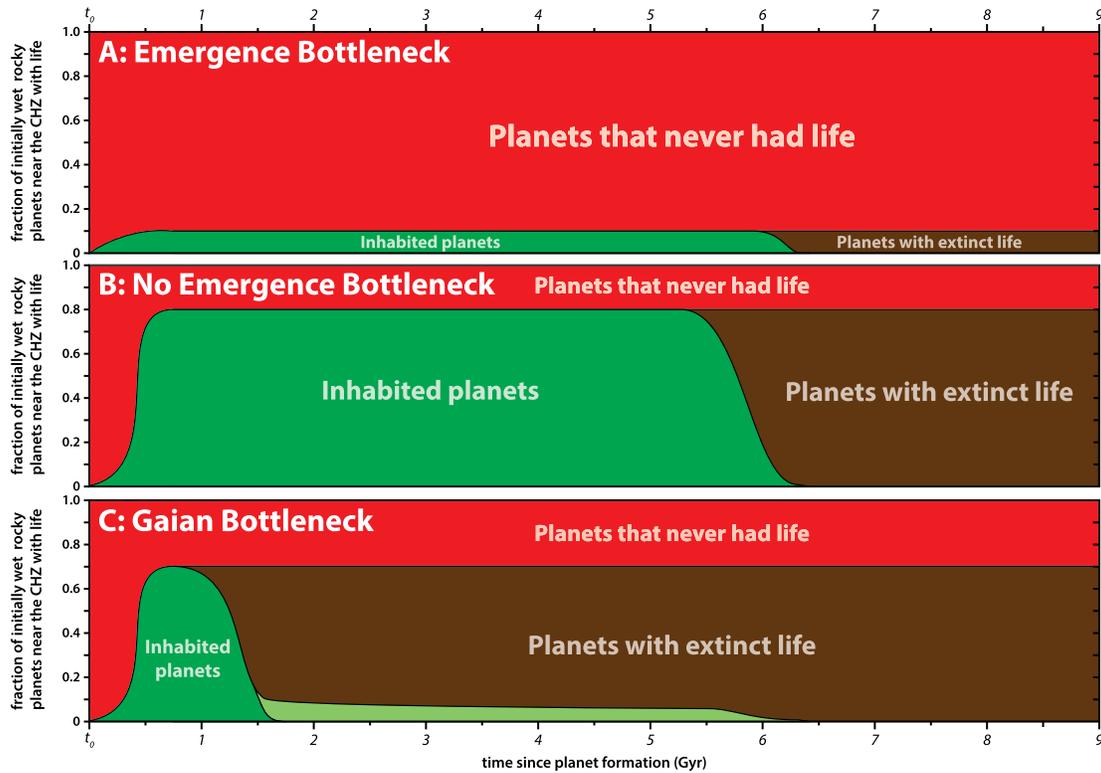


FIG. 7. Different bottleneck scenarios and their fossil predictions. **(A)** Emergence Bottleneck. Life rarely emerges even on wet rocky planets. Few planets will have life or even fossils of extinct life. On the few planets where life does emerge, it persists for billions of years. **(B)** No Emergence Bottleneck. Life emerges with high probability and usually persists for billions of years. Thus, life will be abundant on planets throughout the Universe. There will be many planets where life persisted for billions of years and then went extinct. On the oldest uninhabited planets, fossils of complex life will be abundant. **(C)** Gaian Bottleneck. Life emerges with some probability (possibly quite high), but it goes extinct within a billion years (green). Alternatively, some small fraction of inhabited planets successfully pass through the Gaian bottleneck (light green). The Gaian bottleneck model predicts that the vast majority of the fossils in the Universe will be from extinct microbial life. (Color graphics available at www.liebertonline.com/ast)

to evolve complex and perhaps multicellular forms. In panel C, which illustrates the Gaian bottleneck model, almost all emerging life goes extinct rapidly and therefore, does not have time to evolve into more complex forms. However, even planets with Gaian regulation will not be able to counter indefinitely the increasing luminosity due to stellar evolution (Caldeira and Kasting, 1992; Franck, 2000; Lenton and von Bloh, 2001; Franck *et al.*, 2002; von Bloh *et al.*, 2005). Hence, the extinction at ~ 6 Gyr in all three panels.

If we are able to find well-preserved, ~ 3.8 to ~ 4.3 billion-year-old rocks on Venus or Mars, then we may be able to identify isotopic anomalies produced by biotic actions, in a way analogous to how $^{12}\text{C}/^{13}\text{C}$ ratios are used to infer the existence of the earliest life in Isua, Greenland (Ohtomo *et al.*, 2014). Whether it evolved independently of life on Earth will be difficult to determine. If we find evidence of extant life on Mars or Venus that had an origin independent of Earth life, then this would be evidence against both the Gaian bottleneck hypothesis and the emergence bottleneck.

The surface temperature and existence of liquid water at, or near, the surface could be predominantly due to Gaian regulation rather than abiotic negative feedback. Liquid water on the surface of the planet (particularly old planets) would then not just be a prerequisite for life but a biosignature (Gorshkov *et al.*, 2004). Existence of liquid water

on the surface of a planet may be a better biosignature than oxygen (Luger and Barnes, 2015). Thus, the measurement of exoplanetary surface temperatures compatible with liquid water could be an important part of future search for extraterrestrial life. Remote detection of atmospheric chemical equilibrium may soon develop into a mature science of remote biodetection (*e.g.*, Lovelock and Kaplan, 1975; Krissansen-Totton *et al.*, 2016). The Gaian bottleneck model predicts that the vast majority of the atmospheres of old terrestrial planets in the traditional abiotic CHZs of their host stars will be in chemical equilibrium because they are uninhabited. Hence, atmospheres in chemical disequilibrium will be rare except for young ($t \lesssim 2$ Gyr) terrestrial planets.

In a critique of Gaian logic, Dawkins (1982, p. 236) wrote:

For the analogy [of the Earth as an organism] to apply strictly, there would have to have been a set of rival Gaias, presumably on different planets. Biospheres which did not develop efficient homeostatic regulation of their planetary atmospheres tended to go extinct. The Universe would have to be full of dead planets whose homeostatic regulation systems had failed, with, dotted around, a handful of successful, well-regulated planets of which the Earth is one.

What Dawkins describes here is also a prediction of the Gaian bottleneck hypothesis. The Gaian bottleneck model

parallels evolution on Earth in that the vast majority (~99.9%) of species that have ever lived are now extinct; the vast majority of planetary life has gone extinct.

In the far future, we may be able to find evidence for biogenic isotopic anomalies on the initially wet rocky planets around most stars. Since life does not persist for long in the Gaian bottleneck model, it predicts a universe filled with isotopic or microscopic fossils from the kind of life that can evolve in ~1 billion years, not the fossils of larger multicellular eukaryotes or anything else that would take several billion years to evolve.

Cockell (2014) divided all environments in the Universe into three types: (1) uninhabitable, (2) uninhabited habitats, or (3) inhabited habitats (Cockell, 2011; Zuluaga *et al.*, 2014). A prediction of the Gaian bottleneck (persistence-is-hard) model is that (2) and (3) will be rare. This is unfortunate for future colonization efforts since uninhabited, but habitable, planets are the most ethically appealing places—autochthonous life would not have to be displaced.

Our search for life beyond Earth may be thwarted by the short timescales over which planets may remain inhabited. If it takes several billion years to develop radio telescopes, then the Gaian bottleneck ensures that the vast majority of life in the Universe is either young and microbial, or extinct. Therefore, the Gaian bottleneck model is consistent with current SETI results and can help resolve Fermi's paradox, although it is not one of the solutions to the paradox listed by Webb (2002).

7. What Could Be Wrong with Our Argument?

- (1) Gaian regulation is a controversial idea. It is usually invoked to explain the long-term stability of the surface temperature of Earth, starting in the Proterozoic (~2.5 Gya). So invoking early, pre-Proterozoic Gaia is even more controversial.
- (2) If estimates of sub-aerial continental crust (*e.g.*, by Flament *et al.*, 2008) are significantly too low and there was abundant sub-aerial crust earlier than ~3.8 Gya (*e.g.*, Van Kranendonk, 2010), then abiotic negative feedback based on the carbonate-silicate cycle could have stabilized surface temperatures very early in Earth's history, without Gaian regulation. If early continent formation is a common feature of rocky planets, then invoking Gaian regulation may be unnecessary to explain Earth's early thermal stability.
- (3) We are arguing that
 - (a) Gaian regulation evolved on Earth.
 - (b) The evolution of Gaian regulation is not common.

This could seem paradoxical, because to justify (a) we have presented arguments making the emergence of Gaian regulation plausible. These same arguments could suggest that Gaian regulation would be common. However, there is a large class of phenomena that did happen on Earth that are uncommon or nonexistent elsewhere. We can trace the evolution of these phenomena and explain how they evolved on Earth, but these explanations cannot be generalized. For example, the evolution of the English language can be traced and understood and made plausible, but this plausibility cannot be turned into a generic argument for the evolution of English on other planets.

Humanlike intelligence may be another member of this class (Lineweaver, 2008). We are suggesting that the evolution of a kind of life that can quickly and effectively regulate the volatiles on the surface of the planet on which it finds itself may be a quirky product of the biological evolution of life on early Earth.

- (4) Why should Gaian regulation be rare? Franck *et al.* (2002, 2001) concluded that there are ~500,000 "sister Gaias" in the Milky Way Galaxy but then listed a number of physical factors [*e.g.*, those discussed by Ward and Brownlee (2000)] that could reduce this estimate.
- (5) It could be that early heavy impacts almost always extinguish life and do not need any help from subsequent volatile evolution away from habitability. In which case, Gaian regulation would have nothing to do with life's early or late extinction. Evidence for this would be an anomalously low impact rate on Earth (compared to the early impact rates on other wet rocky planets). If this were the case, there would be no tendency for life to evolve and regulate its environment. In this context, Tyrrell (2013) suggested that "Lucky Gaia" as described by Free and Barton (2007) and the anthropic principle provide a better explanation for the continued habitability of Earth than "Probable Gaia," which we assume here.

If the probability of the emergence and persistence of life on wet rocky planets were infinitesimally small and there were only one life-harboring planet in the Universe, we would, of necessity, find ourselves on that planet. Therefore, our mere existence cannot be used to infer the probability of life elsewhere. Additionally, invoking a scenario in which persistent life on Earth is exceptional compared to other planets (as we do here) cannot be criticized with an argument such as, if Gaian regulation is rare, then we should not be here. Self-selection overcomes this critique.

- (6) One argument against early Gaian regulation is that the unit of biological selection starts small and moves to larger groups as in the following chronological sequence: genes, chromosomes, single cells, colonies of cells (bacterial mats), multicellular organisms, colonies of multicellular organisms (superorganisms), ecosystems of various sizes (which produce Gaian regulation only when they are widespread). In this sequence, the evolution of Gaian regulation happens last. However, among the earliest life-forms we know of, stromatolites were already ecosystems of bacterial mats (Walter *et al.*, 1992).
- (7) The Universe does not seem to be teeming with life. This could be an observational selection effect: it is teeming with life, but we just have not been able to detect it yet. Even in the future when the remote detection of biosignatures is possible, it will be difficult to detect subsurface life (Boston *et al.*, 1992; Gaidos *et al.*, 1999; Jones *et al.*, 2011; McMahon *et al.*, 2013) that does not interact with the planet's atmosphere and is completely sustained by free energy based on geochemical disequilibrium. This would undermine some of the motivation for the Gaian bottleneck model but not the other arguments presented here.

8. Conclusion

We are proposing a potentially universal sequence of events on initially wet rocky planets that can be summarized thusly:

First ~0.5 Gyr: Hot, high bombardment, uninhabitable.

~0.5 to ~1.0 Gyr: Cooler, reduced bombardment, continuous volatile loss.

~0.5 to ~1.0 Gyr: Emergence of life in an environment with a tendency to evolve away from habitability.

~1.0 to ~1.5 Gyr: Inability to maintain habitability, followed by extinction. As a rare alternative, this period would experience the rapid evolution of Gaian regulation and the maintenance of habitability, followed by the persistence of life for several billion more years.

Between the early heat pulses, freezing, volatile content variation, and runaway positive feedbacks, maintaining life on an initially wet rocky planet in the habitable zone may be like trying to ride a wild bull. Most life falls off. Life may be rare in the Universe, not because it is difficult to get started, but because habitable environments are difficult to maintain during the first billion years.

In the book *Vital Dust*, de Duve (1995) presented the case that water and energy are common and abiogenesis may be a cosmic imperative. The most important constraint on the existence of life in the Universe may be whether life, after emerging and evolving into a biosphere, can evolve global mechanisms rapidly enough to mediate the positive and negative feedbacks of abiotic atmospheric evolution. We hypothesize that the early evolution of biologically mediated negative feedback processes, or Gaian regulation as proposed by Lovelock and Margulis (1974), may be necessary to maintain habitability because of the strength, rapidity, and universality of abiotic positive feedbacks on the surfaces of rocky planets in traditional CHZs.

We argue that the habitable surface environments of rocky planets usually become uninhabitable due to abiotic runaway positive feedback mechanisms involving surface temperature, albedo, and the loss of atmospheric volatiles. Because of the strength, rapidity, and universality of abiotic positive feedbacks in the atmospheres of rocky planets in traditional CHZs, biotic negative feedback or Gaian regulation may be necessary to maintain habitability.

The evolution of biospheric regulation of surface volatiles, temperature, and albedo can become a Gaian bottleneck to the persistence of life. This Gaian bottleneck may be a better explanation for the nonprevalence of life than the traditional emergence bottleneck paradigm.

Acknowledgments

A.C. acknowledges PhD scholarship funding from the Australian Postgraduate Award and the Research School of Earth Sciences at the Australian National University. We thank Jim Kasting, David Catling, and Vickie Bennett for helpful discussions. We thank Norm Sleep and an anonymous reviewer for their insightful feedback and constructive suggestions.

References

Abbot, D.S., Cowan, N.B., and Ciesla, F.J. (2012) Indication of insensitivity of planetary weathering behavior and habitable

- zone to surface land fraction. *Astrophys J* 756, doi:10.1088/0004-637X/756/2/178.
- Abe, Y., Abe-Ouchi, A., Sleep, N.H., and Zahnle, K.J. (2011) Habitable zone limits for dry planets. *Astrobiology* 11:443–460.
- Abramov, O. and Mojzsis, S. (2009) Microbial habitability of the Hadean Earth during the Late Heavy Bombardment. *Nature* 459:419–422.
- Abramov, O., Kring, D.A., and Mojzsis, S.J. (2013) The impact environment of the Hadean Earth. *Chemie der Erde Geochemistry* 73:227–248.
- Annala, A. and Annala, E. (2008) Why did life emerge? *International Journal of Astrobiology* 7:293–300.
- Benner, S.A. (2013) Keynote: planets, minerals and life's origin. *Mineral Mag* 77:686.
- Bland, P.A. (2005) The impact rate on Earth. *Philos Transact A Math Phys Eng Sci* 363:2793–2810.
- Boston, P.J., Ivanov, M.V., and McKay, C.P. (1992) On the possibility of chemosynthetic ecosystems in subsurface habitats on Mars. *Icarus* 95:300–308.
- Bostrom, N. (2008) Where are they? Why I hope the search for extraterrestrial life finds nothing. *MIT Technology Review* May/June:72–77.
- Bovaird, T. and Lineweaver, C.H. (2013) Exoplanet predictions based on the generalized Titius-Bode relation. *Mon Not R Astron Soc* 435:1126–1138.
- Bovaird, T., Lineweaver, C.H., and Jacobsen, S.K. (2015) Using the inclinations of Kepler systems to prioritize new Titius-Bode-based exoplanet predictions. *Mon Not R Astron Soc* 448:3608–3627.
- Branscomb, E. and Russell, M.J. (2013) Turnstiles and bifurcators: the disequilibrium converting engines that put metabolism on the road. *Biochim Biophys Acta Bioenergetics* 1827:62–78.
- Budyko, M.I. (1969) The effect of solar radiation variations on the climate of the Earth. *Tellus A* 21:611–619.
- Burke, C.J., Christiansen, J.L., Mullally, F., Seader, S., Huber, D., Rowe, J.F., Coughlin, J.L., Thompson, S.E., Catanzarite, J., Clarke, B.D., Morton, T.D., Caldwell, D.A., Bryson, S.T., Haas, M.R., Batalha, N.M., Jenkins, J.M., Tenenbaum, P., Twicken, J.D., Li, J., Quintana, E., Barclay, T., Henze, C.E., Borucki, W.J., Howell, S.B., and Still, M. (2015) Terrestrial planet occurrence rates for the Kepler GK dwarf sample. *Astrophys J* 809, doi:10.1088/0004-637X/809/1/8.
- Caldeira, K. and Kasting, J.F. (1992) The life span of the biosphere revisited. *Nature* 360:721–723.
- Catling, D.C., Zahnle, K.J., and McKay, C. (2001) Biogenic methane, hydrogen escape, and the irreversible oxidation of early Earth. *Science* 293:839–843.
- Chambers, J. (2014) Planet formation. In *Treatise on Geochemistry*, Vol. 2, 2nd ed., Elsevier, Amsterdam, pp 55–72.
- Chopra, A., Lineweaver, C.H., Brocks, J.J., and Ireland, T.R. (2010) Palaeoecophylostoichiometrics searching for the elemental composition of the last universal common ancestor. In *Proceedings from 9th Australian Space Science Conference*, National Space Society of Australia, Sydney, Australia.
- Ćirković, M.M. (2009) Fermi's paradox—the last challenge for Copernicanism? *Serbian Astronomical Journal* 178:1–20.
- Cockell, C.S. (2011) Vacant habitats in the Universe. *Trends Ecol Evol* 26:73–80.
- Cockell, C.S. (2014) Types of habitat in the Universe. *International Journal of Astrobiology* 13:158–164.
- Conrad, P.G. and Nealson, K.H. (2001) A non-Earthcentric approach to life detection. *Astrobiology* 1:15–24.

- Darwin, C. (1871) *The Descent of Man, and Selection in Relation to Sex*, John Murray, London.
- Davies, P.C.W. and Lineweaver, C.H. (2005) Finding a second sample of life on Earth. *Astrobiology* 5:154–163.
- Dawkins, R. (1982) *The Extended Phenotype: The Gene as the Unit of Selection*, Oxford University Press, Oxford, UK.
- de Duve, C. (1995) *Vital Dust: Life as a Cosmic Imperative*, Basic Books, New York.
- Dhuime, B., Wuestefeld, A., and Hawkesworth, C.J. (2015) Emergence of modern continental crust about 3 billion years ago. *Nat Geosci* 8:552–555.
- Drake, M.J. (2005) Origin of water in the terrestrial planets. *Meteorit Planet Sci* 40:519–527.
- Drake, M.J. and Righter, K. (2002) Determining the composition of the Earth. *Nature* 416:39–44.
- Driscoll, P. and Bercovici, D. (2013) Divergent evolution of Earth and Venus: influence of degassing, tectonics, and magnetic fields. *Icarus* 226:1447–1464.
- Ehrenfreund, P. and Charnley, S.B. (2000) Organic molecules in the interstellar medium, comets, and meteorites: a voyage from dark clouds to the early Earth. *Annu Rev Astron Astrophys* 38:427–483.
- Eigen, M. and Winkler, R. (1992) *Steps towards Life: A Perspective on Evolution*, Oxford University Press, Oxford, UK.
- Elkins-Tanton, L.T. (2008) Linked magma ocean solidification and atmospheric growth for Earth and Mars. *Earth Planet Sci Lett* 271:181–191.
- Elkins-Tanton, L.T. (2011) Formation of early water oceans on rocky planets. *Astrophys Space Sci* 332:359–364.
- Elkins-Tanton, L.T. (2012) Magma oceans in the inner Solar System. *Annu Rev Earth Planet Sci* 40:113–139.
- England, J.L. (2013) Statistical physics of self-replication. *J Chem Phys* 139:1–8.
- Eschenmoser, A. and Volkan Kisakürek, M. (1996) Chemistry and the origin of life. *Helv Chim Acta* 79:1249–1259.
- Falkowski, P.G., Fenchel, T., and Delong, E.F. (2008) The microbial engines that drive Earth’s biogeochemical cycles. *Science* 320:1034–1039.
- Feulner, G. (2012) The faint young Sun problem. *Rev Geophys* 50:1–29.
- Flament, N., Coltice, N., and Rey, P.F. (2008) A case for late-Archaeon continental emergence from thermal evolution models and hypsometry. *Earth Planet Sci Lett* 275:326–336.
- Franck, S. (2000) Habitable zone for Earth-like planets in the Solar System. *Planet Space Sci* 48:1099–1105.
- Franck, S., Block, A., Bloh, W., Bounama, C., Garrido, I., and Schellnhuber, H.J. (2001) Planetary habitability: is Earth commonplace in the Milky Way? *Naturwissenschaften* 88:416–426.
- Franck, S., von Bloh, W., Bounama, C., Steffen, M., Schönberner, D., and Schellnhuber, H.-J. (2002) Habitable zones and the number of Gaia’s sisters. In *The Evolving Sun and Its Influence on Planetary Environments*, Vol. 269, edited by B. Montesinos, A. Gimenez, and E. F. Guinan, Astronomical Society of the Pacific, San Francisco.
- Free, A. and Barton, N.H. (2007) Do evolution and ecology need the Gaia hypothesis? *Trends Ecol Evol* 22:611–619.
- Fressin, F., Torres, G., Charbonneau, D., Bryson, S.T., Christiansen, J., Dressing, C.D., Jenkins, J.M., Walkowicz, L.M., and Batalha, N.M. (2013) The false positive rate of Kepler and the occurrence of planets. *Astrophys J* 766, doi:10.1088/0004-637X/766/2/81.
- Gaidos, E.J., Neelson, K.H., and Kirschvink, J.L. (1999) Life in ice-covered oceans. *Science* 284:1631–1633.
- Gaidos, E.J., Deschenes, B., Dundon, L., Fagan, K., Menviel-Hessler, L., Moskovitz, N., and Workman, M. (2005) Beyond the principle of plentitude: a review of terrestrial planet habitability. *Astrobiology* 5:100–126.
- Goldblatt, C. (2015) Habitability of waterworlds: runaway greenhouses, atmospheric expansion, and multiple climate states of pure water atmospheres. *Astrobiology* 15:362–370.
- Goldblatt, C. and Zahnle, K.J. (2011) Clouds and the faint young Sun paradox. *Climate of the Past* 7:203–220.
- Goldblatt, C., Claire, M.W., Lenton, T.M., Matthews, A.J., Watson, A.J., and Zahnle, K.J. (2009) Nitrogen-enhanced greenhouse warming on early Earth. *Nat Geosci* 2:891–896.
- Gonzalez, G. (2005) Habitable zones in the Universe. *Orig Life Evol Biosph* 35:555–606.
- Gorshkov, V.G., Makarieva, A.M., and Gorshkov, V.V. (2004) Revising the fundamentals of ecological knowledge: the biota-environment interaction. *Ecological Complexity* 1:17–36.
- Gray, R.H. (2015) The Fermi paradox is neither Fermi’s nor a paradox. *Astrobiology* 15:195–199.
- Güdel, M., Dvorak, R., Erkaev, N., Kasting, J.F., Khodachenko, M.L., Lammer, H., Pilat-Lohinger, E., Rauer, H., Ribas, I., and Wood, B.E. (2014) Astrophysical conditions for planetary habitability. In *Protostars and Planets VI*, edited by H. Beuther, R.S. Klessen, C.P. Dullemond, and T. Henning, University of Arizona Press, Tucson, pp 883–906.
- Hanson, R. (1998, September 15) *The Great Filter—Are We Almost Past It?* <http://mason.gmu.edu/~rhanson/greatfilter.html> or <https://archive.is/J02C9> (accessed January 1, 2016).
- Haqq-Misra, J.D., Domagal-Goldman, S.D., Kasting, P.J., and Kasting, J.F. (2008) A revised, hazy methane greenhouse for the Archean Earth. *Astrobiology* 8:1127–1137.
- Harding, S. and Margulis, L. (2010) Water Gaia: 3.5 thousand million years of wetness on planet Earth. In *Gaia in Turmoil: Climate Change, Biodepletion, and Earth Ethics in an Age of Crisis*, edited by E. Crist and H.B. Rinker, MIT Press, Cambridge, MA, pp 41–60.
- Hardy, A., Caceres, C., Schreiber, M.R., Cieza, L., Alexander, R.D., Canovas, H., and Williams, J.P. (2015) Probing the final stages of protoplanetary disk evolution with ALMA. *Astron Astrophys* 583, doi:10.1051/0004-6361/201526504.
- Hart, M.H. (1979) Habitable zones about main sequence stars. *Icarus* 37:351–357.
- Herbst, E. and van Dishoeck, E.F. (2009) Complex organic interstellar molecules. *Annu Rev Astron Astrophys* 47:427–480.
- Hoffman, P.F. (1998) A Neoproterozoic snowball Earth. *Science* 281:1342–1346.
- Höning, D., Hansen-Goos, H., Airo, A., and Spohn, T. (2014) Biotic vs abiotic Earth: a model for mantle hydration and continental coverage. *Planet Space Sci* 98:5–13.
- Howard, A.W., Marcy, G.W., Bryson, S.T., Jenkins, J.M., Rowe, J.F., Batalha, N.M., Borucki, W.J., Koch, D.G., Dunham, E.W., Gautier, T.N., III, Cleve, J.V., Cochran, W.D., Latham, D.W., Lissauer, J.J., Torres, G., Brown, T.M., Gilliland, R.L., Buchhave, L.A., Caldwell, D.A., Christensen-Dalsgaard, J.R., Ciardi, D., Fressin, F., Haas, M.R., Howell, S.B., Kjeldsen, H., Seager, S., Rogers, L., Sasselov, D.D., Steffen, J.H., Basri, G.S., Charbonneau, D., Christiansen, J., Clarke, B., Dupree, A., Fabrycky, D.C., Fischer, D., Ford, E.B., Fortney, J.J., Tarter, J.C., Girouard, F.R., Holman, M., Johnson, J.A., Klaus, T.C., Machalek, P., Moorhead, A.V., Morehead, R.C., Ragozzine,

- D., Tenenbaum, P., Twicken, J.D., Quinn, S.N., Isaacson, H., Shporer, A., Lucas, P.W., Walkowicz, L.M., Welsh, W.F., Boss, A.P., Devore, E., Gould, A., Smith, J.C., Morris, R.L., Prsa, A., Morton, T.D., Still, M., Thompson, S.E., Mullally, F., Endl, M., and MacQueen, P.J. (2012) Planet occurrence within 0.25 AU of solar-type stars from Kepler. *Astrophys J Suppl Ser* 201, doi:10.1088/0067-0049/2012/15.
- Ingersoll, A.P. (1969) The runaway greenhouse: a history of water on Venus. *Journal of the Atmospheric Sciences* 26:1191–1198.
- Johnson, B. and Goldblatt, C. (2015) The nitrogen budget of Earth. *Earth-Science Reviews* 148:150–173.
- Jones, E.G., Lineweaver, C.H., and Clarke, J.D. (2011) An extensive phase space for the potential martian biosphere. *Astrobiology* 11:1017–1033.
- Kasting, J.F. (1988) Runaway and moist greenhouse atmospheres and the evolution of Earth and Venus. *Icarus* 74:472–494.
- Kasting, J.F. (2012) *How to Find a Habitable Planet*, Princeton University Press, Princeton, NJ.
- Kasting, J.F., Whitmire, D.P., and Reynolds, R.T. (1993) Habitable zones around main sequence stars. *Icarus* 101:108–128.
- Kleidon, A. (2012) How does the Earth system generate and maintain thermodynamic disequilibrium and what does it imply for the future of the planet? *Philos Transact A Math Phys Eng Sci* 370:1012–1040.
- Köberl, C. (2006) The record of impact processes on the early Earth: a review of the first 2.5 billion years. In *Processes on the Early Earth*, Geological Society of America Special Paper 405, edited by W.U. Reimold and R.L. Gibson, Geological Society of America, Boulder, CO, pp 1–22.
- Kopp, R.E., Kirschvink, J.L., Hilburn, I.A., and Nash, C.Z. (2005) The Paleoproterozoic snowball Earth: a climate disaster triggered by the evolution of oxygenic photosynthesis. *Proc Natl Acad Sci USA* 102:11131–11136.
- Kopparapu, R.K., Ramirez, R., Kasting, J.F., Eymet, V., Robinson, T.D., Mahadevan, S., Terrien, R.C., Domagal-Goldman, S., Meadows, V., and Deshpande, R. (2013) Habitable zones around main-sequence stars: new estimates. *Astrophys J* 765, doi:10.1088/0004-637X/765/2/131.
- Krissansen-Totton, J., Bergsman, D.S., and Catling, D.C. (2016) On detecting biospheres from thermodynamic disequilibrium in planetary atmospheres. *Astrobiology* 16:39–67.
- Kulikov, Y.N., Lammer, H., Lichtenegger, H.I.M., Penz, T., Breuer, D., Spohn, T., Lundin, R., and Biernat, H.K. (2007) A comparative study of the influence of the active young Sun on the early atmospheres of Earth, Venus, and Mars. *Space Sci Rev* 129:207–243.
- Kump, L.R., Kasting, J.F., and Crane, R.G. (2009) *The Earth System*, 3rd ed., Pearson Higher Education, Upper Saddle River, NJ.
- Kurokawa, H., Sato, M., Ushioda, M., Matsuyama, T., Moriwaki, R., Dohm, J.M., and Usui, T. (2014) Evolution of water reservoirs on Mars: constraints from hydrogen isotopes in martian meteorites. *Earth Planet Sci Lett* 394:179–185.
- Lammer, H., Bredehöft, J.H., Coustenis, A., Khodachenko, M.L., Kaltenecker, L., Grasset, O., Prieur, D., Raulin, F., Ehrenfreund, P., Yamauchi, M., Wahlund, J.-E., Grießmeier, J.-M., Stangl, G., Cockell, C.S., Kulikov, Y.N., Grenfell, J.L., and Rauer, H. (2009) What makes a planet habitable? *Astronomy and Astrophysics Review* 17:181–249.
- Leconte, J., Forget, F., Charnay, B., Wordsworth, R., and Pottier, A. (2013) Increased insolation threshold for runaway greenhouse processes on Earth-like planets. *Nature* 504:268–271.
- Léger, A., Selsis, F., Sotin, C., Guillot, T., Despois, D., Lammer, H., Ollivier, M., and Brachet, F. (2003) A new family of planets? “Ocean-planets.” In *Proceedings of the Conference on Towards Other Earths: DARWIN/TPF and the Search for Extrasolar Terrestrial Planets*, edited by M. Fridlund and T. Henning, compiled by H. Lacoste, ESA SP-539, ESA Publications Division, Noordwijk, the Netherlands, pp 253–259.
- Lenton, T.M. (1998) Gaia and natural selection. *Nature* 394:439–447.
- Lenton, T.M. and von Bloh, W. (2001) Biotic feedback extends the life span of the biosphere. *Geophys Res Lett* 28:1715–1718.
- Lenton, T.M., Schellnhuber, H.J., and Szathmáry, E. (2004) Climbing the co-evolution ladder. *Nature* 431:913.
- Lineweaver, C.H. (2008) Paleontological tests: human-like intelligence is not a convergent feature of evolution. In *From Fossils to Astrobiology (Records of Life on Earth and the Search for Extraterrestrial Biosignatures)*, Vol. 12, edited by J. Seckbach and M. Walsh, Springer, Dordrecht, the Netherlands, pp 353–368.
- Lineweaver, C.H., and Chopra, A. (2012a) The habitability of our Earth and other Earths: astrophysical, geochemical, geophysical, and biological limits on planet habitability. *Annu Rev Earth Planet Sci* 40:597–623.
- Lineweaver, C.H. and Chopra, A. (2012b) What can life on Earth tell us about life in the Universe? In *Genesis—In The Beginning: Precursors of Life, Chemical Models and Early Biological Evolution*, edited by J. Seckbach, Springer, Dordrecht, the Netherlands, pp 799–815.
- Lineweaver, C.H. and Egan, C.A. (2008) Life, gravity and the second law of thermodynamics. *Phys Life Rev* 5:225–242.
- Lodders, K., Palme, H., and Gail, H.-P. (2009) Abundances of the elements in the Solar System. In *Landolt-Börnstein, New Series, Astronomy and Astrophysics*, Vol. VI/4B, Chapter 4.4, edited by J. Trümper, Springer-Verlag, Berlin, pp 560–630.
- Lovelock, J.E. (2000) *Gaia: A New Look at Life on Earth*, Oxford University Press, Oxford, UK.
- Lovelock, J.E. (2005) *Gaia: Medicine for an Ailing Planet*, Gaia Books, London.
- Lovelock, J.E. and Kaplan, I.R. (1975) Thermodynamics and the recognition of alien biospheres [and discussion]. *Philos Trans R Soc Lond B Biol Sci* 189:167–181.
- Lovelock, J.E. and Margulis, L. (1974) Atmospheric homeostasis by and for the biosphere: the Gaia hypothesis. *Tellus*, 26(1–2).
- Lovelock, J.E. and Whitfield, M. (1982) Life span of the biosphere. *Nature* 296:561–563.
- Luger, R. and Barnes, R. (2015) Extreme water loss and abiotic O₂ buildup on planets throughout the habitable zones of M dwarfs. *Astrobiology* 15:1–26.
- Maher, K.A. and Stevenson, D.J. (1988) Impact frustration of the origin of life. *Nature* 331:612–614.
- Marchi, S., Bottke, W.F., Elkins-Tanton, L.T., Bierhaus, M., Wuennemann, K., Morbidelli, A., and Kring, D.A. (2014) Widespread mixing and burial of Earth’s Hadean crust by asteroid impacts. *Nature* 511:578–582.
- Marcy, G.W., Weiss, L.M., Petigura, E.A., Isaacson, H., Howard, A.W., and Buchhave, L.A. (2014) Occurrence and core-envelope structure of 1–4 Earth-size planets around Sun-like stars. *Proc Natl Acad Sci USA* 111:12655–12660.

- Martin, W.F. and Russell, M.J. (2007) On the origin of biochemistry at an alkaline hydrothermal vent. *Philos Trans R Soc Lond B Biol Sci* 362:1887–1925.
- Mat, W.-K., Xue, H., and Wong, J.T.-F. (2008) The genomics of LUCA. *Front Biosci* 13:5605–5613.
- McMahon, S., O’Malley-James, J., and Parnell, J. (2013) Circumstellar habitable zones for deep terrestrial biospheres. *Planet Space Sci* 85:312–318.
- Morbidelli, A., Chambers, J., Lunine, J.I., Petit, J.M., Robert, F., Valsecchi, G.B., and Cyr, K.E. (2000) Source regions and timescales for the delivery of water to the Earth. *Meteorit Planet Sci* 35:1309–1320.
- Morbidelli, A., Lunine, J.I., O’Brien, D.P., Raymond, S.N., and Walsh, K.J. (2012) Building terrestrial planets. *Annu Rev Earth Planet Sci* 40:251–275.
- Nealson, K.H. and Conrad, P.G. (1999) Life: past, present and future. *Philos Trans R Soc Lond B Biol Sci* 354:1923–1939.
- Nealson, K.H. and Rye, R. (2013) Evolution of metabolism. In *Treatise on Geochemistry*, Vol. 10, 2nd ed., edited by H. Palme and H. O’Neill, Elsevier, Amsterdam, pp 43–59.
- Nisbet, E.G. (2002) Fermor Lecture: The influence of life on the face of the Earth: garnets and moving continents. *Geol Soc Spec Publ* 199:275–307.
- Nisbet, E.G. and Fowler, C.M.R. (2014) The early history of life. In *Treatise on Geochemistry*, Vol. 10, 2nd ed., edited by H. Palme and H. O’Neill, Elsevier, Amsterdam, pp 1–42.
- Nisbet, E.G. and Sleep, N.H. (2001) The habitat and nature of early life. *Nature* 409:1083–1091.
- Nisbet, E.G., Zahnle, K.J., Gerasimov, M.V., Helbert, J., Jaumann, R., Hofmann, B.A., Benzerara, K., and Westall, F. (2007) Creating habitable zones, at all scales, from planets to mud micro-habitats, on Earth and on Mars. *Space Sci Rev* 129:79–121.
- Nisbet, E.G., Fowler, C., and Nisbet, R.E.R. (2012) The regulation of the air: a hypothesis. *Solid Earth* 3:87–96.
- Ohtomo, Y., Kakegawa, T., Ishida, A., Nagase, T., and Rosing, M.T. (2014) Evidence for biogenic graphite in early Archaean Isua metasedimentary rocks. *Nat Geosci* 7:25–28.
- O’Malley-James, J.T., Greaves, J.S., Raven, J.A., and Cockell, C.S. (2015) In search of future Earths: assessing the possibility of finding Earth analogues in the later stages of their habitable lifetimes. *Astrobiology* 15:400–411.
- Orgel, L. (1998) The origin of life—a review of facts and speculations. *Trends Biochem Sci* 23:491–495.
- Pace, N.R. (2001) The universal nature of biochemistry. *Proc Natl Acad Sci USA* 98:805–808.
- Pavlov, A.A., Hurtgen, M.T., Kasting, J.F., and Arthur, M.A. (2003) Methane-rich Proterozoic atmosphere? *Geology* 31:87–90.
- Petigura, E.A., Howard, A.W., and Marcy, G.W. (2013) Prevalence of Earth-size planets orbiting Sun-like stars. *Proc Natl Acad Sci USA* 110:19273–19278.
- Pope, E.C., Bird, D.K., and Rosing, M.T. (2012) Isotope composition and volume of Earth’s early oceans. *Proc Natl Acad Sci USA* 109:4371–4376.
- Raymond, S.N., Quinn, T., and Lunine, J.I. (2004) Making other Earths: dynamical simulations of terrestrial planet formation and water delivery. *Icarus* 168:1–17.
- Raymond, S.N., Quinn, T., and Lunine, J.I. (2007) High-resolution simulations of the final assembly of Earth-like planets. 2. Water delivery and planetary habitability. *Astrobiology* 7:66–84.
- Raymond, S.N., O’Brien, D.P., Morbidelli, A., and Kaib, N.A. (2009) Building the terrestrial planets: constrained accretion in the inner Solar System. *Icarus* 203:644–662.
- Ricklefs, R.E. and Miller, G.L. (2000) Homeostasis depends upon negative feedback. In *Ecology*, 4th ed., edited by R.E. Ricklefs and G.L. Miller, W.H. Freeman, New York, pp 92–93.
- Robles, J., Lineweaver, C.H., Grether, D., Flynn, C., Egan, C.A., Pracy, M.B., Holmberg, J., and Gardner, E. (2008) A comprehensive comparison of the Sun to other stars: searching for self-selection effects. *Astrophys J* 684:691–706.
- Rosing, M.T., Bird, D.K., Sleep, N.H., Glassley, W., and Albarède, F. (2006) The rise of continents—an essay on the geologic consequences of photosynthesis. *Palaeogeogr Palaeoclimatol Palaeoecol* 232:99–113.
- Rosing, M.T., Bird, D.K., Sleep, N.H., and Bjerrum, C.J. (2010) No climate paradox under the faint early Sun. *Nature* 464:744–747.
- Russell, M.J., Nitschke, W., and Branscomb, E. (2013) The inevitable journey to being. *Philos Trans R Soc Lond B Biol Sci* 368:1–19.
- Sagan, C. and Chyba, C.F. (1997) The early faint Sun paradox: organic shielding of ultraviolet-labile greenhouse gases. *Science* 276:1217–1221.
- Sagan, C. and Mullen, G. (1972) Earth and Mars: evolution of atmospheres and surface temperatures. *Science* 177:52–56.
- Schneider, S.H. (2004) *Scientists Debate Gaia: The Next Century*, MIT Press, Cambridge, MA.
- Schneider, S.H. and Boston, P.J. (1991) *Scientists on Gaia*, MIT Press, Cambridge, MA.
- Schwartzman, D. and Volk, T. (1989) Biotic enhancement of weathering and the habitability of Earth. *Nature* 340:457–460.
- Shields, A.L., Barnes, R., Agol, E., Charnay, B., Bitz, C., and Meadows, V.S. (2016) The effect of orbital configuration on the possible climates and habitability of Kepler-62f. *Astrobiology*, in press, doi:10.1089/ast.2015.1353.
- Sleep, N.H., Zahnle, K.J., Kasting, J.F., and Morowitz, H. (1989) Annihilation of ecosystems by large asteroid impacts on the early Earth. *Nature* 342:139–142.
- Sleep, N.H., Zahnle, K.J., and Lupu, R.E. (2014) Terrestrial aftermath of the Moon-forming impact. *Philos Transact A Math Phys Eng Sci* 372, doi:10.1098/rsta.2013.0172.
- Smith, E. and Morowitz, H.J. (2016) *The Origin and Nature of Life on Earth: The Emergence of the Fourth Geosphere*, Cambridge University Press, Cambridge, UK.
- Sousa, F.L. and Martin, W.F. (2014) Biochemical fossils of the ancient transition from geoenergetics to bioenergetics in prokaryotic one carbon compound metabolism. *Biochim Biophys Acta Bioenergetics* 1837:964–981.
- Stüeken, E.E., Anderson, R.E., Bowman, J.S., Brazelton, W., Colangelo-Lillis, J., Goldman, A.D., Som, S.M., and Baross, J.A. (2013) Did life originate from a global chemical reactor? *Geobiology* 11:101–126.
- Tarter, J.C. (2001) The search for extra-terrestrial intelligence. *Annu Rev Astron Astrophys* 39:511–548.
- Tian, F. (2015) Atmospheric escape from Solar System terrestrial planets and exoplanets. *Annu Rev Earth Planet Sci* 43:459–476.
- Tielens, A.G.G.M. (2013) The molecular universe. *Rev Mod Phys* 85:1021–1081.
- Tyrrell, T. (2013) *On Gaia: A Critical Investigation of the Relationship between Life and Earth*, Princeton University Press, Princeton, NJ.

- Van Kranendonk, M.J. (2010) Two types of Archean continental crust: plume and plate tectonics on early Earth. *Am J Sci* 310:1187–1209.
- Villanueva, G.L., Mumma, M.J., Novak, R.E., Hartogh, P., Encrenaz, T., Tokunaga, A., Khayat, A., and Smith, M.D. (2015) Strong water isotopic anomalies in the martian atmosphere: probing current and ancient reservoirs. *Science* 348:218–221.
- von Bloh, W., Bounama, C., and Franck, S. (2005) Dynamic habitability of extrasolar planetary systems. In *A Comparison of the Dynamical Evolution of Planetary Systems*, edited by R. Dvorak and S. Ferraz-Mello, Springer, Dordrecht, the Netherlands, pp 287–300.
- Walker, J.C.G. (1985) Carbon dioxide on the early Earth. *Orig Life Evol Biosph* 16:117–127.
- Walker, J.C.G., Hays, P.B., and Kasting, J.F. (1981) A negative feedback mechanism for the long-term stabilization of Earth's surface temperature. *J Geophys Res Oceans* 86:9776–9782.
- Walter, M., Grotzinger, J.P., and Schopf, J.W. (1992) Proterozoic stromatolites. In *The Proterozoic Biosphere: A Multidisciplinary Study*, edited by J.W. Schopf, C. Klein, and D.J. Des Marais, Cambridge University Press, Cambridge, UK, pp 253–260.
- Ward, P.D. and Brownlee, D. (2000) *Rare Earth: Why Complex Life is Uncommon in the Universe*, Copernicus Books, New York.
- Watson, A.J. and Lovelock, J.E. (1983) Biological homeostasis of the global environment: the parable of Daisyworld. *Tellus B* 35B:284–289.
- Webb, S. (2002) *If the Universe Is Teeming with Aliens... Where Is Everybody?: Fifty Solutions to the Fermi Paradox and the Problem of Extraterrestrial Life*, Copernicus Series, Springer Science and Business Media, New York.
- Zuluaga, J.I., Salazar, J.F., Cuartas-Restrepo, P., and Poveda, G. (2014) The habitable zone of inhabited planets. *Biogeosci Discuss* 11:8443–8483.

Address correspondence to:
 Aditya Chopra
 Research School of Earth Sciences
 The Australian National University
 142 Mills Road
 Acton, ACT, 2601
 Australia
 E-mail: aditya.chopra@anu.edu.au

Submitted 22 September 2015
 Accepted 10 October 2015

Abbreviations Used

AHZ = abiogenesis habitable zone
 CHZ = circumstellar habitable zone